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# ORIGINAL ARTICLE

# **EURO-CORDEX:** new high-resolution climate change projections for European impact research

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**Abstract** A new high-resolution regional climate change ensemble has been established for Europe within the World Climate Research Program Coordinated Regional Downscaling Experiment (EURO-CORDEX) initiative. The first set of simulations with a horizontal resolution of 12.5 km was completed for the new emission scenarios RCP4.5 and RCP8.5 with more simulations expected to follow. The aim of this paper is to present this data set to the different communities active in regional climate modelling, impact

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#### L. M. Bouwer

Institute for Environmental Studies, Vrije Universiteit, De Boelelaan 1087, 1081 HV Amsterdam, The Netherlands assessment and adaptation. The EURO-CORDEX ensemble results have been compared to the SRES A1B simulation results achieved within the ENSEMBLES project. The large-scale patterns of changes in mean temperature and precipitation are similar in all three scenarios, but they differ in regional details, which can partly be related to the higher resolution in EURO-CORDEX. The results strengthen those obtained in ENSEMBLES, but need further investigations. The analysis of impact indices shows that for RCP8.5, there is a substantially larger change projected for temperature-based indices than for RCP4.5. The difference is less pronounced for precipitation-based

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Laboratoire de Météorologie Dynamique (IPSL), Ecole Polytechnique, 91128 Palaiseau Cedex, France indices. Two effects of the increased resolution can be regarded as an added value of regional climate simulations. Regional climate model simulations provide higher daily precipitation intensities, which are completely missing in the global climate model simulations, and they provide a significantly different climate change of daily precipitation intensities resulting in a smoother shift from weak to moderate and high intensities.

**Keywords** Regional climate change · Impact indices · EURO-CORDEX · Heat wave · Heavy precipitation · Dry spells

## Introduction

Climate impact assessments and the development of regional to local-scale adaptation strategies require the availability of high-resolution climate change scenarios, including an assessment of their robustness and their inherent uncertainties. The WCRP Coordinated Regional Downscaling Experiment (CORDEX, http://wcrp-cordex.ipsl.jussieu.fr/; Giorgi et al. 2006) provides an internationally coordinated framework to improve regional climate scenarios. This includes harmonisation of model evaluation activities in the individual modelling centres and the generation of multi-model ensembles of regional climate projections for the land-regions worldwide.

As part of the global CORDEX framework, the EURO-CORDEX initiative (http://www.euro-cordex.net/) provides regional climate projections for Europe at 50 km (EUR-44) and 12.5 km (EUR-11) resolution, thereby complementing coarser resolution data sets of former activities like, e.g., PRUDENCE and ENSEMBLES. The regional simulations are downscaling the new CMIP5

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David Kinloch Michie Chair of Rural Economy & Environmental Sustainability, Institute of Geography & the Lived Environment, School of GeoSciences, The University of Edinburgh, Drummond Street, Edinburgh EH8 9XP, UK global climate projections (Taylor et al. 2012) and the new representative concentration pathways (RCPs) (Moss et al. 2010; van Vuuren et al. 2011). Twenty-six modelling groups contributing 11 different regional climate models, partly in different model configurations, actively support EURO-CORDEX.

In its initial phase, EURO-CORDEX mainly focussed on model evaluation in present-day climate (e.g., Vautard et al. 2013; Kotlarski et al. 2013). So far more than 30 evaluation simulations have been conducted. Further activities include the coordinated analysis of future climate simulations, the joint analysis of dynamical and empiricalstatistical methods and the design of suitable bias correction techniques to tailor EURO-CORDEX data for direct application in climate impact research. Particular emphasis is put on the construction of a simulation matrix that covers uncertainty in emission scenarios, the driving global climate model and the downscaling method in the best affordable manner. Here, we present the first results of the high-resolution (EUR-11) future climate simulations from EURO-CORDEX.

We used for comparison the results obtained from the regional climate projections in the FP6 ENSEMBLES project (Hewitt and Griggs 2004; van der Linden and Mitchell 2009). A number of regional limited-area models were used to downscale transient global climate projections over Europe at a 25 and 50 km resolution over the second half of the twentieth century and along the twenty-first century. The ENSEMBLES climate projections were carried out under the assumptions of the Special Report on Emission Scenario (SRES) A1B scenario (IPCC 2000). This scenario follows the storyline of the IPCC family of A1 scenarios. It assumes a rapid economic growth and development of technologies, with a worldwide population peaking in the middle of the twenty-first century, and a

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R. Vautard · P. Yiou Laboratoire des Sciences du Climat et de l'Environnement, IPSL, CEA/CNRS/UVSQ, Gif sur Yvette, France balanced use of energy resources. This scenario leads to a rapid increase in fossil CO2 emissions until 2050 and a decrease afterwards. As compared to other SRES scenarios, the CO2 emissions lie in the middle of the scenario range.

EURO-CORDEX scenario simulations use the new Representative Concentration Pathways (RCPs) defined for the Fifth Assessment Report of the IPCC (Moss et al. 2010). In contrast to the SRES scenarios, RCP scenarios do not specify socioeconomic scenarios, but assume pathways to different target radiative forcing at the end of the twenty-first century. For instance, scenario RCP8.5 assumes an increase in radiative forcing of 8.5 W/m2 by the end of the century relative to pre-industrial conditions.

A comparison between the climate effects of SRES and RCP scenarios (Rogelj et al. 2012) indicates that the A1B scenario leads to a global mean temperature increase in the likely range of 2.8–4.2 °C, which approximates to RCP6 and lies clearly between RCP4.5 and RCP8.5.

The aim of the present study is to present a new highresolution (12.5 km) data set from a multi-model multiscenario ensemble of regional climate simulations for impact research. Such a high resolution has not been reached before in previous climate model projections. The ENSEMBLES project covered RCM simulations for Europe with a maximum resolution of 25 km, and in PRU-DENCE, all simulations were done on a 50 km grid.

The analysis carried out here is directed towards regional climatic changes in Europe, addressing the differences of mean changes in annual mean temperature and total precipitation for the scenarios A1B, RCP4.5 and RCP8.5. In addition to the mean values, a range of climate indices important for climate impact studies in different sectors were calculated for sub-regions of Europe including: heavy precipitation events, dry spells and heat waves.

# Experimental setup, statistical methods and definitions of indices

## Experimental setup

In total, nine simulations have been done for RCP4.5 and ten for RCP8.5. They were all performed on the EURO-CORDEX domain, expect one simulation, which covered only the MED-CORDEX domain. The size and location of the CORDEX domains can be seen on the CORDEX web page (http://wcrp-cordex.ipsl.jussieu.fr/). The EURO-CORDEX domain covers all countries in the European Union, but it does not map perfectly to the Europe region defined for the IPCC Fifth Assessment Report. For the eastern part of Turkey, unfortunately, no regional model projections are available.

A summary of the grid configuration and differences in the parameterisation schemes for the participating regional models (ALADIN5.1: Colin et al. 2010, Herrmann et al. 2011, CCLM: Rockel et al. 2008, HIRHAM: Christensen et al. 1998, RACMO2: Meijgaard van et al. 2012, RCA4: Samuelsson et al. 2011, Kupiainen et al. 2011, REMO: Jacob et al. 2012, WRF Version 3.3.1: Skamarock et al. 2008) is given in the supplementary material (Table s1). The RCP scenarios, the driving GCMs and the driven RCMs as well as the simulation length are listed in the supplementary material (Table s2a). Seven different RCMs and five different GCMs have been used in this study. Two of the RCMs were driven by four/five different GCMs; five GCM-RCM chains did simulate both RCP scenarios. They all provide data at least until the mid of the century. Eight RCP4.5 simulations and nine RCP8.5 simulations had reached the end of the century.

Two additional sensitivity studies have been performed to study the influence of the smaller domain size used by one model (supplementary material, Figure s6) and the effect of using one RCM more often than others (supplementary material, Figure s7).

From the ENSEMBLES data set, 20 transient RCM simulations reaching the end of the century are used for most of the analysis (supplementary material, Table s2b). Maximum and minimum temperature values were only available from nine simulations. The ensemble includes simulations on 50 and 25 km grid scale, but most simulations were carried out on a 25 km grid (16 out of 20 and 7 out of 9). The coarser resolution simulations were included to enlarge the sample size for the statistical analyses. A sensitivity study using only the 25 km simulations for the analyses of changes in the mean fields of temperature and precipitation and in the indices showed that there are only minor differences in the horizontal pattern of the meteorological parameters as well as in the regions with a significant and/or robust change (not shown).

#### Statistical methods

All regional changes have been analysed for the land areas of the EURO-CORDEX domain using daily data for three time periods: 1971–2000, 2021–2050 and 2071–2100. For changes in all outcome measures (mean annual and seasonal temperature, total annual and seasonal precipitation, heavy precipitation, heat waves and dry spells), significance and robustness were tested using a method adapted from Pfeifer et al. (2013). This method identifies regions with relatively strong and robust climate changes from an ensemble of climate change simulations. It can be applied to simulation results on regular model grids or to data aggregated onto larger regions. The robustness of the information given by the ensemble of climate projections is analysed using two statistical tests. The first test regards the agreement of the simulations in terms of the direction of the changes. In the second test, the significance of the projected changes in each simulation is assessed by applying the Mann–Whitney–Wilcoxon test. Only regions, which pass both tests, are identified as regions with robust projected changes.

The stringency of the tests can be adjusted by the choice of several parameters, for example the significance level, or by shifting the percentage of simulations, which have to pass the individual tests. We visualise the regions with relatively strong and robust climate changes. Regions with non-significant changes are treated in the same way as regions where the models strongly disagree on the direction of the changes. It is thus by purpose not possible to distinguish between regions where to a certain confidence no significant changes are projected to occur and regions where the projected changes cannot be given with sufficient confidence. Insignificant changes are not equivalent to the lack of agreement in the projected changes and should be interpreted differently, depending on the context.

First, the direction of change was assessed. Changes in regions in which more than 66 % of the models agree in the direction of change were called robust changes. For the analyses of the significance, two different data preparation methods were used. For total precipitation and maximum, mean and minimum temperature, one value each per year and grid box was calculated (30 values per simulation per grid box). For heat waves, dry spells and heavy precipitation, only one value exists for each simulation and grid box, since they are defined over a 30-year period. Again they were calculated for all three time periods. For significance testing, the Mann–Whitney–Wilcoxon test was applied to the entire ensemble data set for the respective time periods.

For the climate change signal of mean annual temperature and total annual precipitation, a spatial correlation is calculated for the respective parameters of the SRES A1B and RCP8.5 simulation results for the end of the century based on the five sub-regions. For this analysis, all parameters for both scenarios were remapped to the ENSEMBLES grid with 25 km resolution.

# Definitions of impact-relevant indices

In addition to the calculation of mean temperature and precipitation changes, we focussed on high-impact phenomena. Heavy precipitation, dry spells and heat waves were calculated according to the following definitions:

 "Heavy precipitation" is defined as the intensity of the heavy precipitation events defined as the 95th percentile of daily precipitation (only days with precipitation >1 mm/day are considered).

- "Dry spells" are defined as periods of at least 5 consecutive days with daily precipitation below 1 mm. The 95th percentile of the length of all identified dry spells is considered in this analysis and referred to as "extended dry spells."
- For the analysis of the change in number of heat waves, we used two different definitions. In the first definition, heat waves were considered as periods of more than three consecutive days exceeding the 99th percentile of the daily maximum temperature of the May to September season of the control period (1971–2000). The second definition, representing more extreme heat wave events, is based on the WMO definition (Frich et al. 2002). Heat waves were defined as periods of more than 5 consecutive days with daily maximum temperature of the May to September season for the control period (1971–2000) by at least 5 °C.

A range of indices was selected based on known climate impacts in Europe for infrastructure, agriculture and human health (Alcamo et al. 2007). The impact indices are calculated as means for 5 European sub-regions used in the European chapter of the forthcoming IPCC Fifth Assessment Report. The sub-regional classification is based on Metzger et al. (2005) (supplementary material, Fig. s1) and selected to represent 5 different climate and environmental characteristics: Alpine, Atlantic, Continental, Northern and Southern Europe. The list of impact index definitions is given in http://cccma.seos.uvic.ca/ETCCDI/list\_27\_indices. shtm.

For all ensembles, the "likely" range is defined as range within the 17th and 83rd percentile of projected changes.

# Results

Mean changes of temperature and precipitation

To date, the regional climate change projections provided by the EU-FP6 ENSEMBLES multi-model ensemble for the SRES A1B scenario are considered as state-of-the-art for European climate impact research. Therefore, we compared this data set to the new regional EURO-COR-DEX data set for RCP8.5 and RCP4.5. Figure 1 shows the ensemble mean of the mean annual temperature and total annual precipitation change until the end of the century. Figure 1b, d shows a robust and statistically significant warming, with regional differences, in the range of 1–4.5 °C for RCP4.5 and of 2.5–5.5 °C for RCP8.5. These ranges encompass the warming range projected for the A1B scenario, where temperature increases between 3 and







A1B

: robust

(e)

70°N

65°N

60°N

55°N

50°N

45°N

40°N

35°N

10°W

0°



15°F

60°E

[K]

5

4

3

2.5

1.5

2

1

.5

4.5

3.5

45°F

RCP85

15°W

(d)

5

-5

30°W







20°E

30°E

10°E

Hatched areas indicate regions with robust and/or statistical significant change (a, c, e). Changes are robust and significant across the entire European continent (b, d, f)

Deringer

**Table 1** Spatial correlation of SRES A1B and RCP8.5 emissionscenarios for changes in mean annual temperature and annual totalprecipitation of the sub-regions for the time periods 2021–2050 and2071–2100

Spatial	Climate para	ameter		
correlation of RCP8.5 and SRES A1B	Mean annua temperature	1	Annual total precipitation	1
	2021-2050	2071-2100	2021-2050	2071-2100
Alpine	0.88	0.95	0.92	0.94
Atlantic	0.82	0.98	0.87	0.94
Continental	0.94	0.96	0.72	0.92
Northern	0.97	0.97	0.59	0.81
Southern	0.90	0.89	0.71	0.96

4.5 °C (Fig. 1f). The projected spatial patterns are very similar in all scenarios with greater annual mean warming in Southern Europe and towards the northeast. Under RCP8.5, large parts of Northern Scandinavia, Eastern Europe and the Alpine ridge might be exposed to a warming of more than 4.5 °C compared to 1971–2000, which could be avoided by RCP4.5.

Associated with the large increase in temperature in RCP8.5 are robust changes in annual precipitation. The ensemble mean projects a statistically significant increase in large parts of Central Europe and Northern Europe of up to about 25 % and a decrease in Southern Europe. A zone with small changes, which are not significant (however, partially robust in the sign of change), indicates where the climate change signals change the sign (Fig. 1c, white areas). The pattern of the changes is very similar for RCP4.5, but less pronounced (Fig. 1a). The spatial pattern for A1B precipitation changes qualitatively agrees with the described changes for RCP4.5 and RCP8.5, and the magnitude of the changes mostly lies in-between the two RCPs. However, differences in the spatial patterns are seen over the British Isles, Benelux and Germany (Fig. 1e).

For mean temperature and precipitation change, a spatial correlation has been done between RCP8.5 and A1B results (Table 1). For all sub-regions, the spatial correlation between SRES A1B and RCP8.5 is very high, with 0.82–0.97 for temperature changes and 0.59–0.92 for precipitation changes depending on the region for mid-century. Towards the end of the century, the correlation is even stronger for both parameters.

Seasonal changes of mean temperature and precipitation are shown in the supplementary material (Fig. s2-s5). The seasonal temperature change signals show more regional heterogeneity than the annual mean. The zone between regions in which precipitation increases in the north and decreases in the south shifts southwards in summer and northwards in winter. Seasonal mean changes for heavy precipitation

The projected seasonal mean changes in heavy precipitation for the three emission scenarios are relatively similar, but some regional differences are visible (Figs. 2 (RCP8.5), 3 (A1B), 4 (RCP4.5)). Most obvious differences are the increased regional detail in the RCP8.5 and RCP4.5, which is related to the higher horizontal resolution of about 12.5 km for the RCPs compared to 25 km for A1B, for which more homogeneous changes are calculated. The annual cycle of changes in heavy precipitation is similar in all three scenarios, but the amplitude of the change is stronger in RCP8.5 than in A1B in several regions. The results for RCP8.5 include a possible decrease in heavy summer precipitation by about 25 % over some parts of the Iberian Peninsula and Southern France, accompanied by regional increases in parts of Spain and Portugal. For winter, RCP8.5 projects strongest increases in heavy precipitation (up to 35 %) in Central and Eastern Europe, whereas A1B projects changes up to 25 % only in this region. Only for some parts in Scandinavia, A1B shows similar values as RCP8.5.

Important regional differences in heavy precipitation are projected for the RCP4.5 scenario. Compared to RCP8.5, the seasonal patterns of change are similar, but the amount of change is much smaller (up to 15 % in large areas with isolated spots up to 25 %) and—besides isolated regions in Southern Europe (mostly along coastlines)—no decrease in heavy precipitation is indicated.

GCM simulations tend to underestimate the high precipitation intensities (Sun et al. 2006). An improved distribution of high precipitation intensities is an important advantage of regional climate simulations. Figure 5 shows the relative frequencies of daily precipitation intensities of an ensemble of five GCM simulations and the corresponding regional downscaling experiments for the reference period 1971-2000 analysed over a central land region of the EURO-CORDEX domain (45°N-50°N and 2°E-17°E). For this analysis, data from all grid cells were taken into account. The distribution illustrates that the GCMs generally produce more precipitation intensities with up to 12 mm/day. The RCMs, in contrast, show higher intensities. Strong intensities above 30 mm/day do hardly occur in the GCM simulations. Figure 6 shows the temporal changes of the precipitation frequencies between the near future period and the reference period for both scenarios and the GCM and the RCM ensembles. Striking is that both model types reduce the number of weak precipitation intensities below 9 mm/day in both scenarios and increase the relative frequencies in all higher intensity classes. This shift in daily precipitation intensities, however, turns out much more moderate in the RCM than in the GCM simulations of both scenarios, the RCP4.5 and the RCP8.5.



Fig. 2 Projected seasonal changes of heavy precipitation (%) based on the RCP8.5 scenario for 2071–2100 compared to 1971–2000. *Hatched areas* indicate regions with robust and/or statistical significant change

The frequency changes of the RCM simulations are especially in the range between 10 and 20 mm/day less than half of the GCM changes. Above 30 mm/day, however, the increase in the RCM ensembles exceeds the climate change signal of the GCMs. Figure 6 also demonstrates that both effects—the reduction in weak intensities and the increase in strong intensities—are more pronounced in the RCP8.5 scenario. The analyses of the processes leading to the different behaviour are beyond the scope of this paper and will be studied in a separate paper.

This analysis proves two effects of an increased resolution, which can be regarded as an added value of regional climate simulations. On the one hand, the RCMs provide higher daily precipitation intensities, which are completely missing in the GCM simulations, and on the other hand, they provide a significantly different climate change of daily precipitation intensities resulting in a smoother shift from weak to moderate and high intensities.

# Mean length of dry spells

Projected changes in the 95th percentile of the mean length of dry spells are shown in Fig. 7 for A1B (top), RCP8.5 (middle) and RCP4.5 (bottom) for 2021–2050 and 2071–2100 with respect to 1971–2000. For the early period, the change patterns are very similar in all scenarios (left row), although the number of simulations taken into account for each scenario ensemble is different. Beside some common features in South-West Europe, substantial differences in the projected changes for dry spells lengths



Fig. 3 Projected seasonal changes of heavy precipitation (%) based on the A1B scenario for the period 2071–2100 compared to 1971–2000. *Hatched areas* indicate regions with robust and/or statistical significant change

are visible until 2100. For RCP4.5 and A1B, a small increase in the length of extended dry spells is projected for Central Europe, which is more pronounced in A1B. A decrease in the length of extended dry spells is calculated in A1B for parts of Scandinavia. This feature is extended towards the Alps in the RCP8.5, in which the number of dry spells increases (not shown). This means that under RCP8.5 more but shorter dry spells are projected in the alpine region. For regions with a large increase in the length of extended dry spells is decreasing (not shown).

# Mean number of heat waves

Projected changes in the mean number of heat waves during May–September are presented in Fig. 8, for RCP4.5 and RCP8.5, for the two future time periods and for two different definitions of heat waves. From the upper four panels, displaying the p99-heat wave definition, it is obvious that with less warming (see "Mean changes of temperature and precipitation" section) in RCP4.5, the increase in number of heat waves is smaller than in RCP8.5. This is more pronounced towards the end of the century (Fig. 8c, d) than for the earlier time period (Fig. 8a, b). For both scenarios, the increase is strongest in Southern Europe, but towards the end of the century the number of heat waves increases all over Europe. The number of heat waves for Southern Europe is projected to increase by more than 45. The increase is mostly robust and significant. The change in the number of heat waves considerably depends on the definition (thresholds and duration), which is used. Therefore, a second definition was



Fig. 4 Projected seasonal changes of heavy precipitation (%) based on the RCP4.5 scenario for the period 2071–2100 compared to 1971–2000. *Hatched areas* indicate regions with robust and/or statistical significant change

used based on that developed by the World Meteorological Organization (see "Definitions of impact-relevant indices" section). Under this definition, the increase in the mean number of heat waves is much less (Fig. 8e, f). For the WMO-heat wave definition not a single heat wave is detected in the ensemble mean for the reference period as well as for mid of the century, because the criteria are much stricter. Also the duration of the heat wave is two days longer than in the p99-heat wave definition. For RCP8.5, meaning under the strongest projected warming, towards the end of the century, an increase is only projected for some parts of Southern Europe with additional 5 to more than 9 heat waves. The increase is significant and robust south of 55° latitude.

Indices by sub-region

Projected changes of several impact indices, which could be of interest for impact studies in different sectors, are listed in Tables 2 (A1B) and 3 (RCP4.5 and RCP8.5) for the 5 sub-regions. For all sub-regions and indices, the median shifts into the same direction, independent of the scenario. For almost all indices, a substantially larger change in the median is projected in RCP8.5, compared to RCP4.5, however, the likely ranges frequently overlap. Exceptions to this are annual total precipitation in the Atlantic and Continental sub-regions, tropical nights in the Northern sub-region and the cold spell duration everywhere. Here the projected changes in RCP4.5 and RCP8.5



Fig. 5 Frequencies of daily precipitation intensities of an ensemble of five GCM simulations and the corresponding RCMs for the reference period 1971–2000 over a central land region of the EURO-CORDEX domain ( $45^{\circ}N$ – $50^{\circ}N$  and  $2^{\circ}E$ – $17^{\circ}E$ )



**Fig. 6** Temporal changes of the precipitation frequencies between the near future period (2021–2050) and the reference period (1971–2000) for RCP4.5 and RCP8.5 and the GCM and the RCM ensembles

are rather similar. Differences between the RCP scenarios are most pronounced for growing season length and warm spell duration index, with no overlap between the likely ranges over all sub-regions (even the full range seldom overlaps).

The median change in A1B is generally centred within RCP4.5 and RCP8.5. For some cases like annual total precipitation in the Southern sub-region or tropical nights in the Atlantic, Northern and Southern sub-regions, however, the median change in A1B is even stronger than the median change in the RCP8.5 and only for change in annual total precipitation in the Continental sub-region, the median change in A1B is lower than the median change in RCP4.5. The spreads of the projected changes defined as the likely ranges are generally the same between RCP4.5 and RCP8.5 or slightly larger in RCP8.5. Exceptions to this are frost days in the Continental sub-region, tropical nights in the Northern sub-region and total rainfall amount above the 99th percentile of daily rain (wet days only) in the Southern sub-region. Here the likely ranges of projected changes are larger in the RCP4.5 scenario.

## **Conclusions and outlook**

Regional climate change patterns for Europe projected by the high-resolution regional climate change ensemble within the EURO-CORDEX initiative for the new emission scenarios RCP4.5 and RCP8.5 have been compared to the state-of-the-art regional climate change data set from ENSEMBLES. The overall spatial patterns for temperature and precipitation changes and related indices are similar. There is a large degree of consistency between the ensembles of the three emission scenarios. The results clearly strengthen the previous findings obtained from the ENSEMBLES data set; however, there are some important new findings.

Climate projections from the new ensemble indicate a reduced northwards shift of Mediterranean drying evolution and slightly stronger mean precipitation increases over most of Europe. This is consistent with the two evaluation papers by Vautard et al. (2013) and Kotlarski et al. (2013), for reanalysis-driven hindcast simulations, where it was found that higher resolution leads to more precipitation in the climate models.

The high-resolution in the EURO-CORDEX simulations is clearly visible in the change pattern for heavy precipitation events. This spatial information is potentially very useful for climate impact studies. The more detailed spatial patterns in the high-resolution simulations can be related to better resolved physical processes like convection and heavy precipitation, and due to better representation of surface characteristics and their spatial variability. This needs to be elaborated further to understand the physical processes involved and the robustness of the pattern.

For some regions, especially the Alps, a connection between the changes in length of extended dry spells and the number of all dry spells (not only the longest ones) is (a)

70°N

65°N

60°N

55°N

50°N

45°N

40°N

35°N

45°N

40°N

35°N

30°N

10°W

10°W

٥٥







45°N

40°N

30°N

10°W

09

: significant
 : robust

10°E

20°E

45°N

40°N 35°N

35°N

Fig. 7 Projected changes in the 95th percentile of the length of dry spells (days) for 2021-2050 compared to 1971-2000 (a, c, e) and 2071-2100 compared to 1971-2000 (b, d, f) for A1B (a, b), RCP8.5

10°E

0 : significant
 : robust 20°E

30°E

(c, d) and RCP4.5 (e, f) scenarios. Hatched areas indicate regions with robust and/or statistical significant changes

30°E

50°N

45°N

40°N



◄ Fig. 8 Projected changes in the mean number of heat waves occurring in the months May–September for 2021–2050 compared to 1971–2000 (a, b) and 2071–2100 compared to 1971–2000 (c through f). Heat waves in figures a through d are defined as periods of more than 3 consecutive days exceeding the 99th percentile of the daily maximum temperature of the May to September season for the control period (1971–2000). Heat waves in figures e and f are defined as periods of more than 5 consecutive days with daily maximum temperature exceeding the mean maximum temperature of the May–September season for the control period (1971–2000) by at least 5 °C. They are based on A1B emission scenario (e), the RCP8.5 (b, d, f) and the RCP4.5 (a, c). *Hatched areas* indicate regions with robust and/or statistical significant change

detected. In this region, the decreasing length of extended dry spells is linked to an increasing number of all dry spells. This seems to be plausible, but follow-up studies are needed to understand the governing processes.

The investigation related to possible changes in heat waves shows clearly that the definition of the heat wave is dominating the regional change pattern more than the differences in scenarios or time periods. Therefore, it is of utmost importance to clearly define the index in the light of the specific study and most likely for the sector of interest.

 Table 2
 Projected changes of selected climate parameters and indices for 2071–2100 with respect to 1971–2000 spatially averaged for European sub-regions for A1B scenario

Scenario A1B	Climate parameters	Measure	Alpine	Atlantic	Continental	Northern	Southern
2071–2100 minus 1971–2000	Mean annual	Median	3.4	2.5	3.3	3.8	3.6
	Temperature in K**	Min	2.8	1.9	2.1	3.2	2.3
		Likely in the range	3.1-4.5	2.1-3.5	2.8-4.5	3.5-5.0	3.3-4.1
		Max	5.4	4.7	5.7	5.8	5.5
	Frost days (1) per	Median	-50	-24	-44	-54	-24
	year*	Min	-37	-13	-26	-38	-12
		Likely in the range	-38 to $-57$	-15 to $-34$	-27 to $-53$	-40 to $-55$	-12 to -31
		Max	-72	-39	-56	-71	-34
	Summer days (2) per	Median	14	21	32	7	48
	year*	Min	4	9	21	3	33
		Likely in the range	11-20	16–32	22-41	5-14	33-51
		Max	21	34	43	27	51
	Tropical nights (4) per	Median	3	8	21	4	47
	year*	Min	1	2	14	1	18
		Likely in the range	2–9	6–17	16–35	1–7	35-52
		Max	11	32	43	10	60
	Growing season	Median	47	41	52	41	36
	length (5) days per	Min	27	23	20	25	14
	growing season**	Likely in the range	34–56	33-51	33-62	27–46	27-41
		Max	75	55	81	61	51
	Warm spell duration	Median	57	44	42	67	91
	index (14) days per	Min	46	29	26	37	67
	year*	Likely in the range	51-84	35-72	37–69	47–96	85-112
		Max	126	125	94	119	144
	Cold spell duration	Median	-5	-5	-6	-6	-5
	index (15) days per	Min	-4	-4	-4	-5	-3
	year*	Likely in the range	-4 to -5	-4 to -6	-5 to -6	-5 to -8	-4 to -5
		Max	-8	-9	-9	-9	-8
	Annual total	Median	7	3	3	16	-15
	precipitation (27)	Min	1	9	-9	4	-7
	ın %**	Likely in the range	5-12	-4 to 5	-1 to 5	13-21	-12 to -18
		Max	15	-11	12	29	-25
	Annual total	Median	57	65	53	64	43
	precipitation where	Min	35	28	31	32	21
	RR > 99p of 1971/2000 (26) in	Likely in the range	47–68	42–98	44–77	47-88	35–57
	%**	Max	117	112	110	105	74

Numbers are based on 9 (indicated with\*) and 20 (indicated with\*\*) regional model simulations. The likely range defines the range of 66 % of all projected changes around the ensemble median

cenarios	Climate parameters	Measure	Alpine		Atlantic		Continental		Northern		Southern	
			RCP 4.5	RCP 8.5	RCP 4.5	RCP 8.5	RCP 4.5	RCP 8.5	RCP 4.5	RCP 8.5	RCP 4.5	RCP 8.5
071-2100 minus 1971-2000	Mean annual tem-perature	Median	2.4	4.6	1.7	3.2	2.1	4.1	2.9	5.2	2.0	4.2
	in K	Min	1,8	3,8	1,3	2,5	1,6	3,6	2,0	4,1	1,9	3,8
		Likely in the range	1.9 - 3.4	3.9-6.0	1.4–2.1	2.7-3.6	1.6 - 3.2	3.7-5.2	2.0-4.2	4.1-6.2	1.9–2.7	3.9-5.4
		Max	3.6	6.3	2.9	4.2	3.2	5.3	4.3	6.5	3.2	5.7
	Frost days (1) per year	Median	-40	-70	-28	-40	-34	-62	-40	-68	-22	-43
		Min	-25	-55	-12	-21	-16	-46	-24	-58	-10	-22
		Likely in the range	-26 to -41	-57 to -85	-15 to $-30$	-26 to -50	-18 to -40	-50 to -65	-26 to -43	-60 to -83	-11 to -30	-23 to -51
		Max	-47	-93	-33	-60	-41	-73	-52	-93	-31	-51
	Summer days (2) per year	Median	8	19	11	24	20	37	4	13	27	54
		Min	3	10	9	17	11	27	2	5	21	43
		Likely in the range	4–14	12–24	6-14	22–28	13-24	30-46	2-16	6-22	25–33	46-60
		Max	18	25	33	38	28	49	23	28	37	67
	Tropical nights (4) per	Median	1	4	3	7	6	22	1	1	20	45
	year	Min	0	1	0	б	2	11	0	0	7	23
		Likely in the range	1–3	2-5	1-5	3-12	9–27	17-31	0-5	1–3	11-24	25-57
		Max	8	6	18	17	30	37	7	13	41	58
	Growing season length (5)	Median	31	61	39	58	26	58	23	55	27	49
	days per growing	Min	23	52	24	41	17	52	17	37	16	34
	20432011	Likely in the range	23–39	52-83	27–43	47–68	20–38	53-71	19–33	41-60	17–33	38–53
		Max	45	95	45	75	41	75	42	78	38	58
	Warm spell duration	Median	34	96	20	65	23	73	35	82	34	124
	index (14) days per year	Min	26	73	17	46	16	52	22	64	28	90
		Likely in the range	29–55	77–136	20-31	49–87	18-42	58-93	23-42	75-113	32–69	98-177
		Max	69	162	55	102	54	106	63	130	83	186
	Cold spell duration index	Median	-5	-5	-5	-5	9-	-0	L	9	-5	-5
	(15) days per year	Min	-3	-3	-2	-4	-3	-5	-4	-4	-3	-4
		Likely in the range	-4 to -7	4 to6	-3 to -6	-4 to -5	-4 to -7	-5 to -8	-6 to -8	-5 to -7	-3 to -5	-4 to -5
		Max	L	-0	-0	-6	L—	-6	-8	-8	-6	9-
	Annual total precipitation	Median	5	14	1	4	6	10	10	22	-0	-10
	(27) in %	Min	3	5		-2	0	0	7	18	-11	0
		Likely in the range	4-8	7–16	-1 to 7	2–9	1-13	4-19	8-17	19–32	-10 to 0	-21 to -2
		Max	12	18	6	6	16	29	22	34	2	-27
	Annual total precipitation	Median	38	79	36	71	44	65	43	82	36	49
	where $RR > 99p$ of 1071 1000 (26) ::. 02	Min	24	41	20	48	17	37	27	64	23	30
	0/ 111 (07) 0007/11/61	Likely in the range	25-60	44–105	25-70	49–105	33-60	44–75	28-65	66-110	31-57	38-58
		Max	73	110	73	118	73	106	02	120	67	65

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For example, the stricter (WMO) definition is more relevant to human health impacts. Any statement about possible changes in heat waves without this detailed information can easily be misleading regarding the risk of increases in future heat waves in Continental and Northern Europe.

The added value of regional climate simulations could be stated. RCMs provide higher daily precipitation intensities than GCMs, and they provide a significantly different climate change of daily precipitation intensities resulting in a smoother shift from weak to moderate and high intensities. Shifts of weak precipitation events are crucial for impact studies, in particular for hydrology, agriculture and air pollution.

The intention for this paper is not an in depth intercomparison of the time periods or scenarios, but the presentation of the new ensemble and a comparative analysis to currently used regional climate change information.

The EURO-CORDEX high-resolution data set will grow continuously and made available to the community. At the time of writing, more ensemble members are being conducted for both, the RCP4.5 and RCP8.5 scenarios. Most simulations will run until 2100 and will enlarge the ensemble. The EURO-CORDEX community agreed on common quality control and will provide a wealth of information for future climate research, impact assessment and adaptation.

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